Atmospheric Dynamics Curiosities

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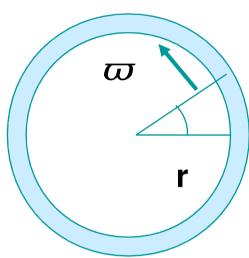
Crew Building room 218

Main structure of course

- Expressing Newton's laws for a fluid on a rotating sphere
- Consequence of scale (Rossby Number) gives geostrophic and higher approximations for wind in terms of pressure (or geopotential)
- Variation of wind with height
 - In free atmosphere "thermal" wind
 - In boundary layer, eddy friction, and Ekmann spiral
- Circulations in the tropics (introduced via "dishpan" results)
- Notions of vorticity, divergence, and potential vorticity
- Use of those concepts in
 - Rossby waves (2-D and 3-D)
 - Cyclone waves

Conservation of angular momentum

View from Pole star



 $\overline{\omega}$ is angular velocity of ring

$$\varpi mr^2 = const$$

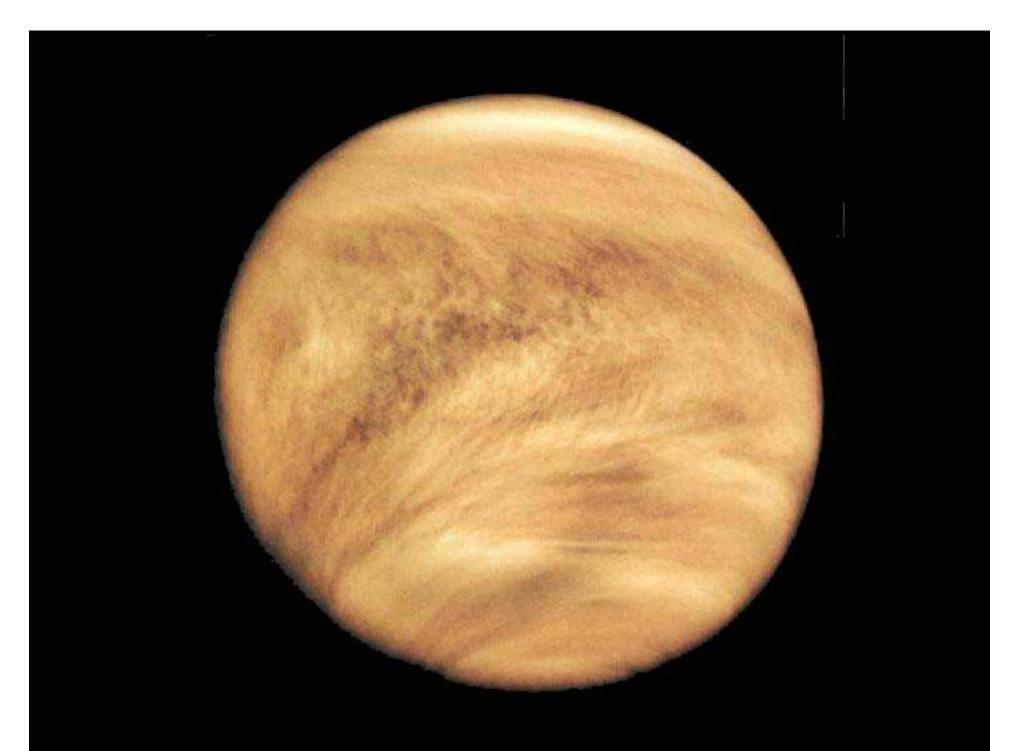
$$\omega r^2 = const$$

$$\varpi = \Omega + \frac{u}{r}$$

$$r = (a+z)\cos\boldsymbol{\varphi}$$

$$\left| \frac{u}{(a+z)\cos \varphi} \right| (a+z)^2 \cos^2 \varphi = const$$

See also Tut. Q 9.3
$$(u+(a+z)\Omega\cos\varphi)(a+z)\cos\varphi = const$$



Venus



- Radius: 6052 km
- Surface Gravity 8.9 m/s²
- Rotation Period: 243 EarthDays (retrograde)
- Orbit Period: 225 Earth Days
- Length of Solar Day: 113 Earth Days
- Surface Temperature: 465°C
- Surface Pressure: 90 bar

Atmospheric Dynamics

DIRECT CIRCULATION:

Superrotation

Atmosphere rotates with a period of 4 days!

From:

Mayr, Harris (1983): Quasiaxisymmetric circulation and superrotation in planetary atmospheres.

Astronomy and Astrophysics 121, 124-136.

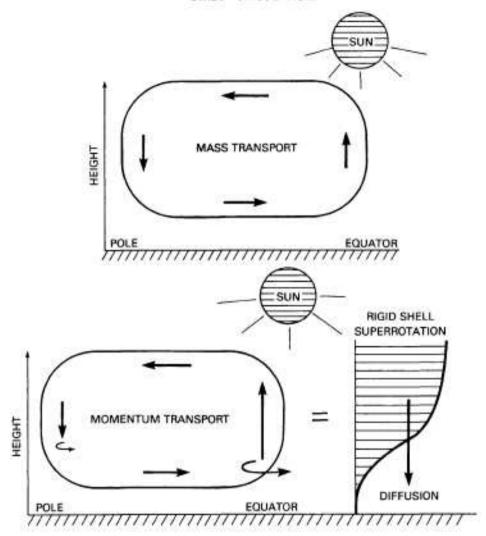
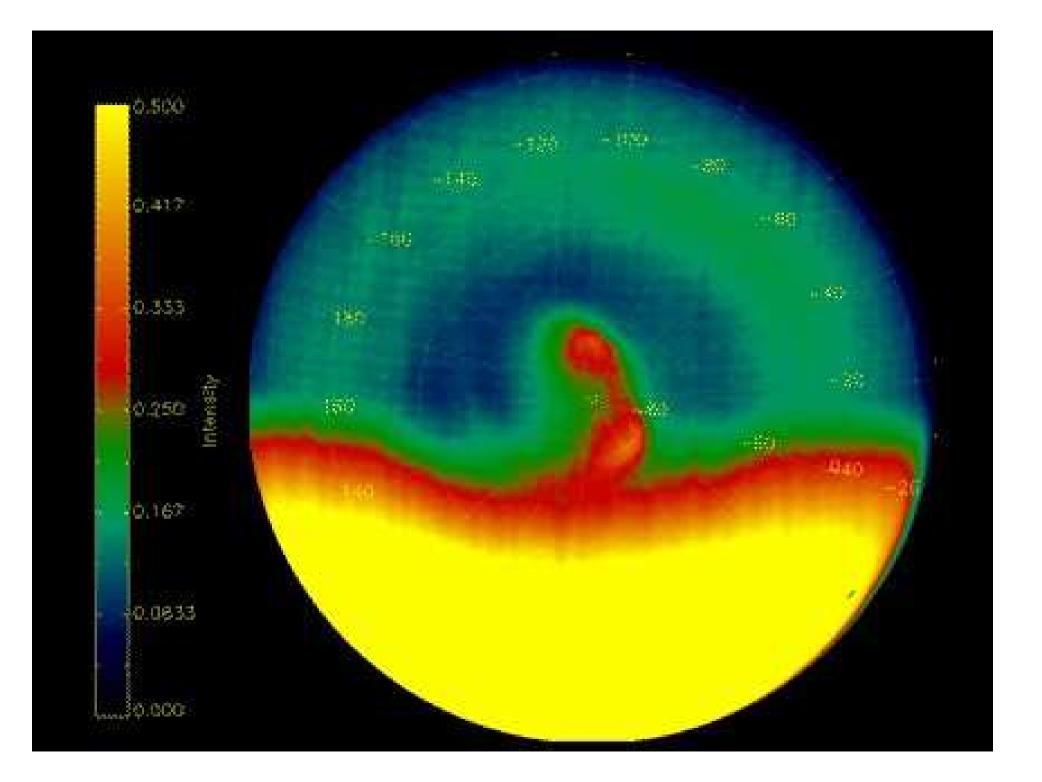
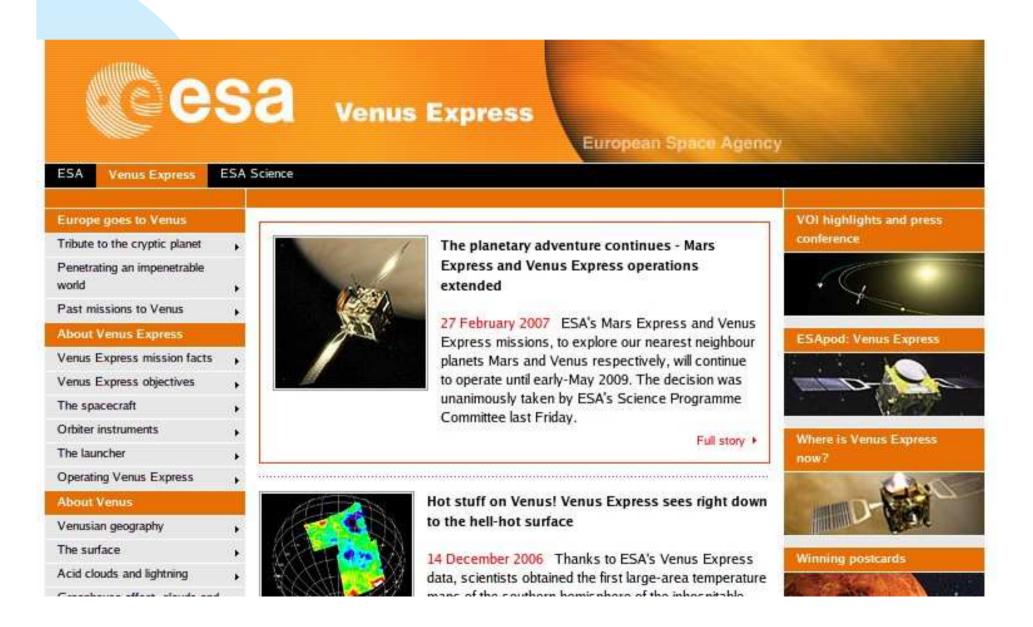


Fig. 1. Schematic illustration of the mass (top)- and momentum (bottom)-budgets for the direct circulation producing the rigid shell component of superrotation which is observed on virtually every planet in our solar system. The balance is established between upward transport of momentum by convection (more being carried upwards at low latitudes than is carried downwards at high latitudes) and diffusion



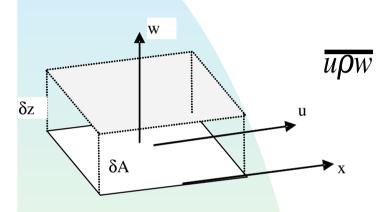
More Info: Venus Express http://www.esa.int/SPECIALS/Venus_Express



Eddy friction

$$\overline{s} = \frac{1}{2Q} \int_{t_o - Q}^{t_o + Q} s dt$$

$$\operatorname{eddy} \quad s' \equiv s - \overline{s}$$



$$= \rho (\overline{u} + u')(\overline{w} + w') = \rho (\overline{u} \overline{w} + u' \overline{w} + \overline{u} w' + u' w')$$

$$= \rho (\overline{u} \cdot \overline{w}) + \rho (\overline{u' w'})$$

Upward flux of x-mom^t due to the eddies =

$$\tau_{xz} = \rho(\overline{u'w'})$$

Accel due to the eddies =
$$-\frac{1}{\rho} \frac{\partial \tau}{\partial z}$$

Introduce eddy diffusion coeff.

$$\tau_{xz} = -\rho K \frac{\partial u}{\partial z}$$

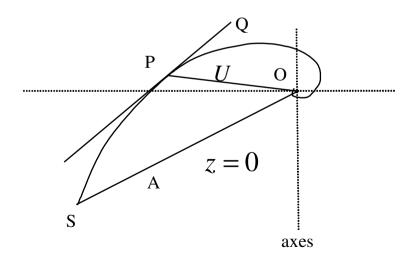
Ekmann Spiral

- Steady unaccelerated flow
- Uniform geostrophic wind in vertical and horizonal
- Variations of density small in bottom km

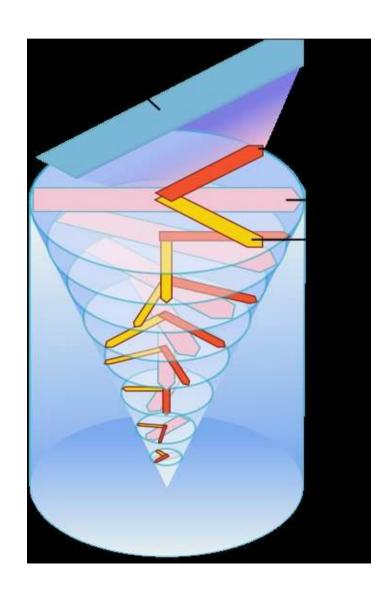
$$fk \wedge \overline{v}_h = -\frac{1}{\rho} \nabla_h p + K \left(\frac{\partial^2 \overline{v}_h}{\partial z^2} \right) \qquad fk \wedge v_a = + K \left(\frac{\partial^2 v_a}{\partial z^2} \right)$$

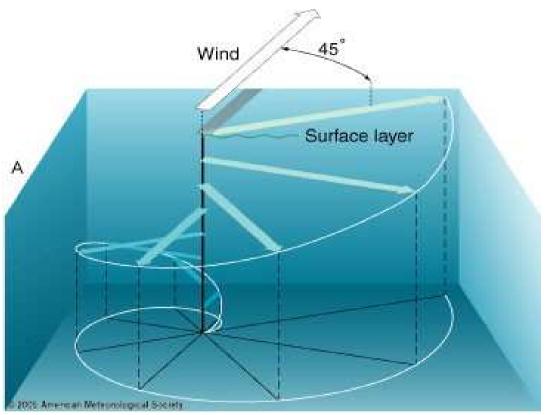
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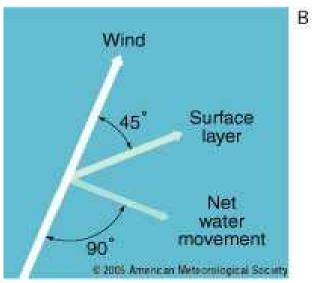
Shape of ageostrophic component



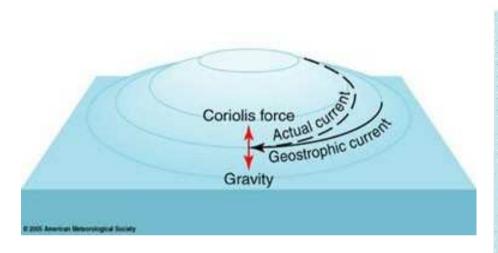
Ekman-Spiral in the Ocean



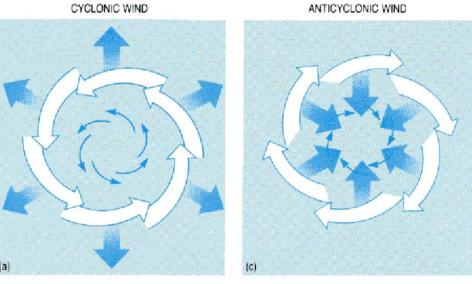


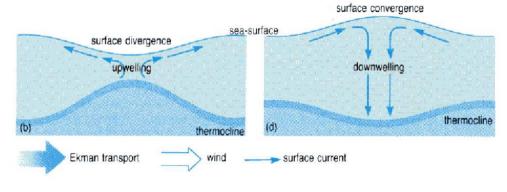


Oceanic Gyres

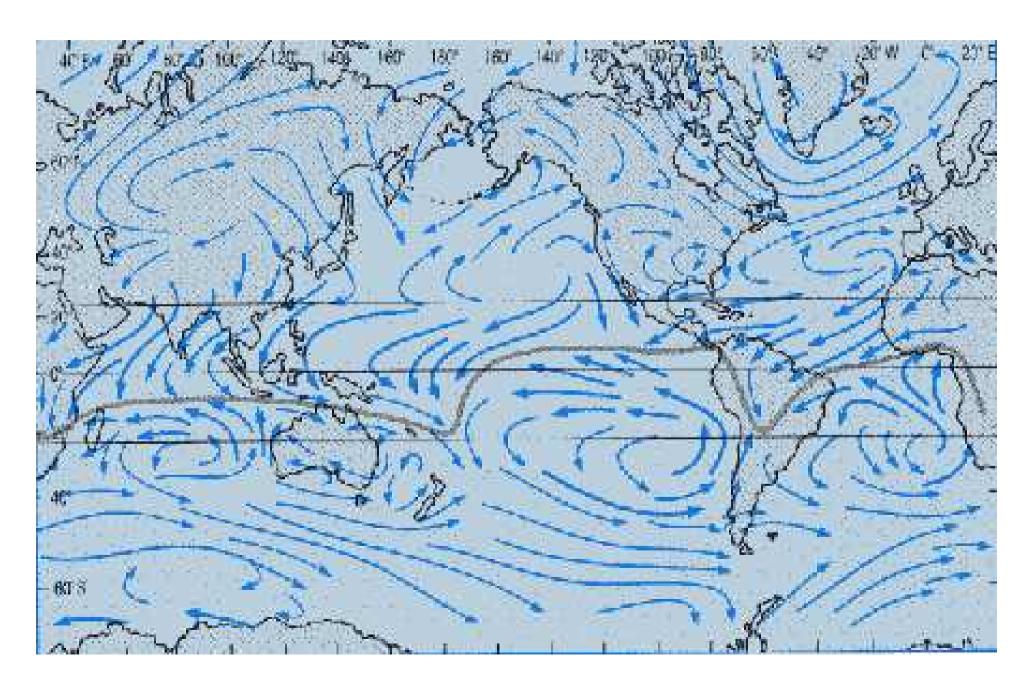


NORTHERN HEMISPHERE

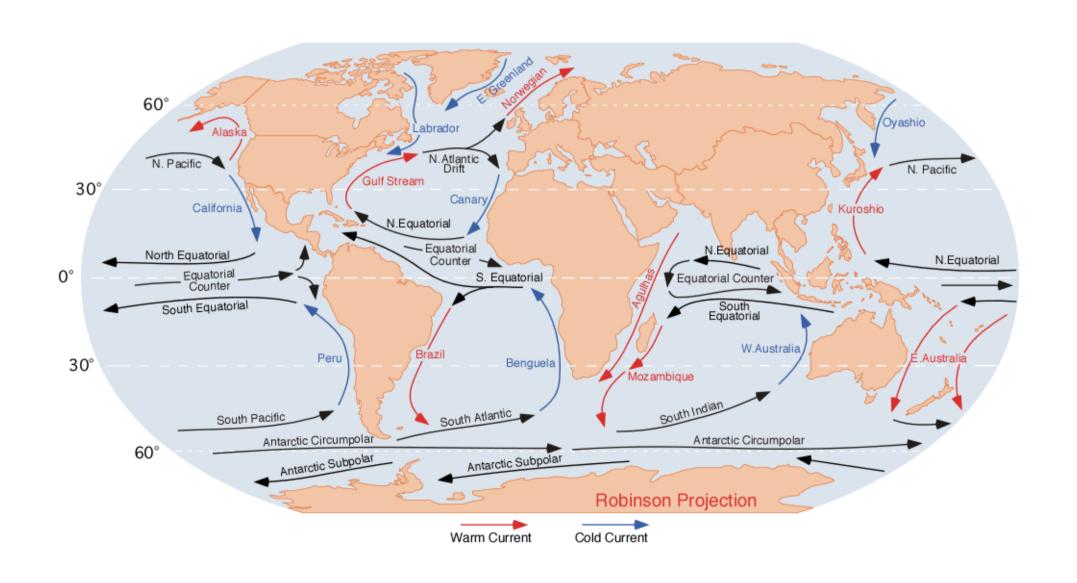




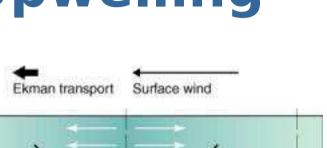
Wind

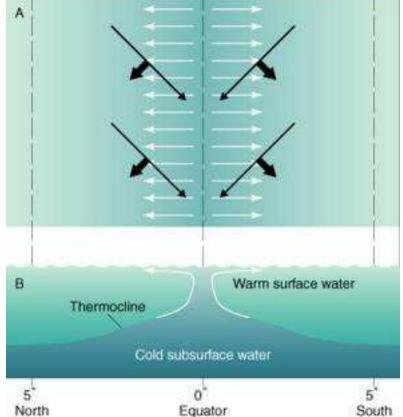


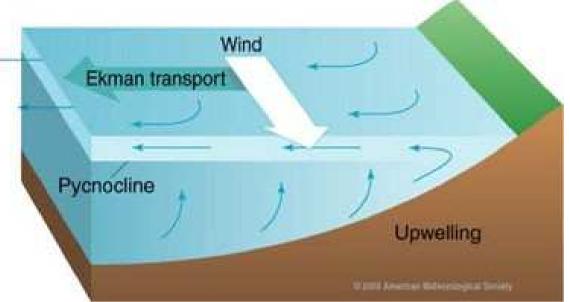
Ocean surface currents

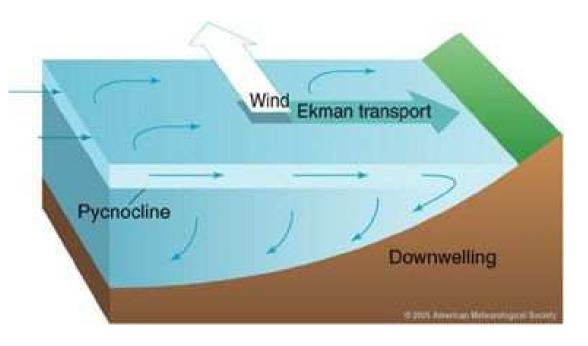


Currents and Upwelling

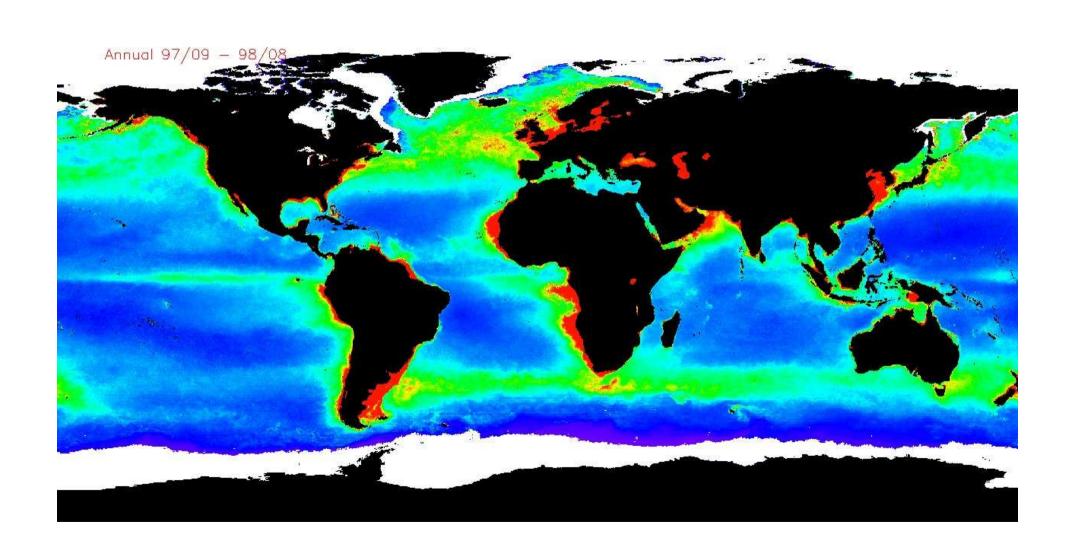




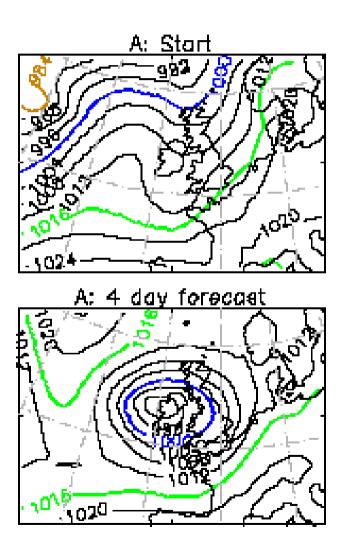


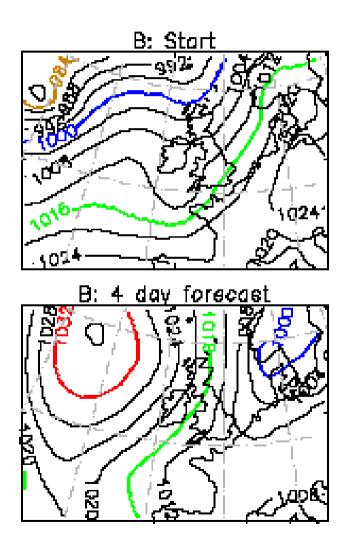


Primary Productivity

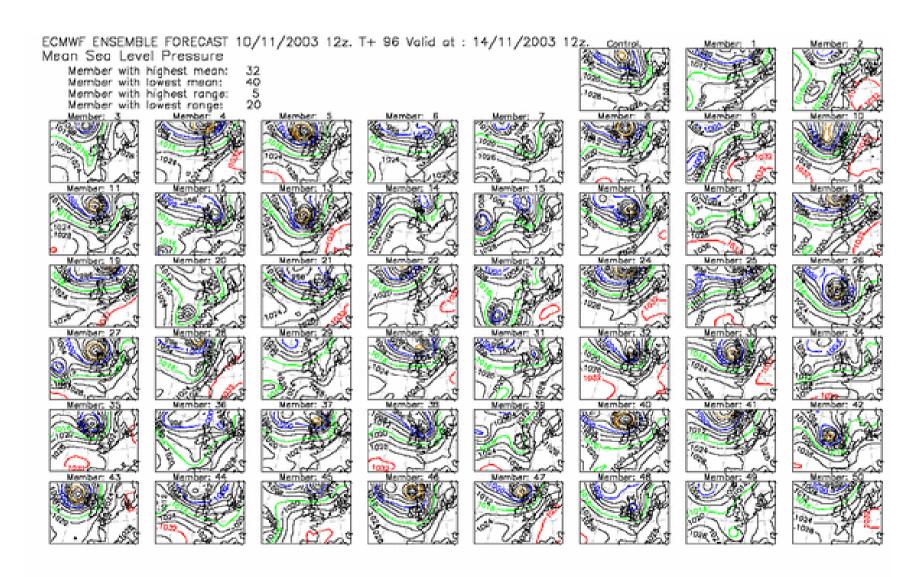


Unpredictable Weather?

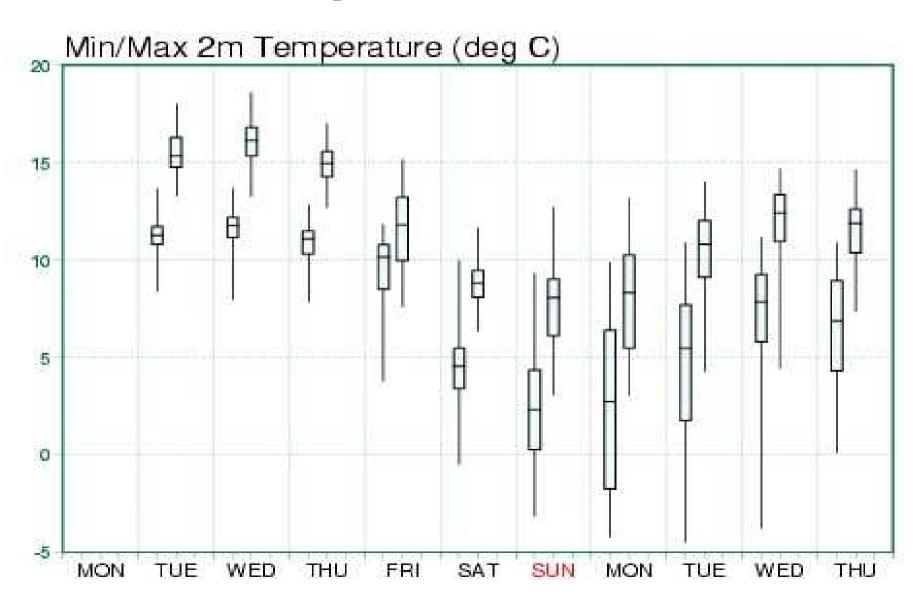


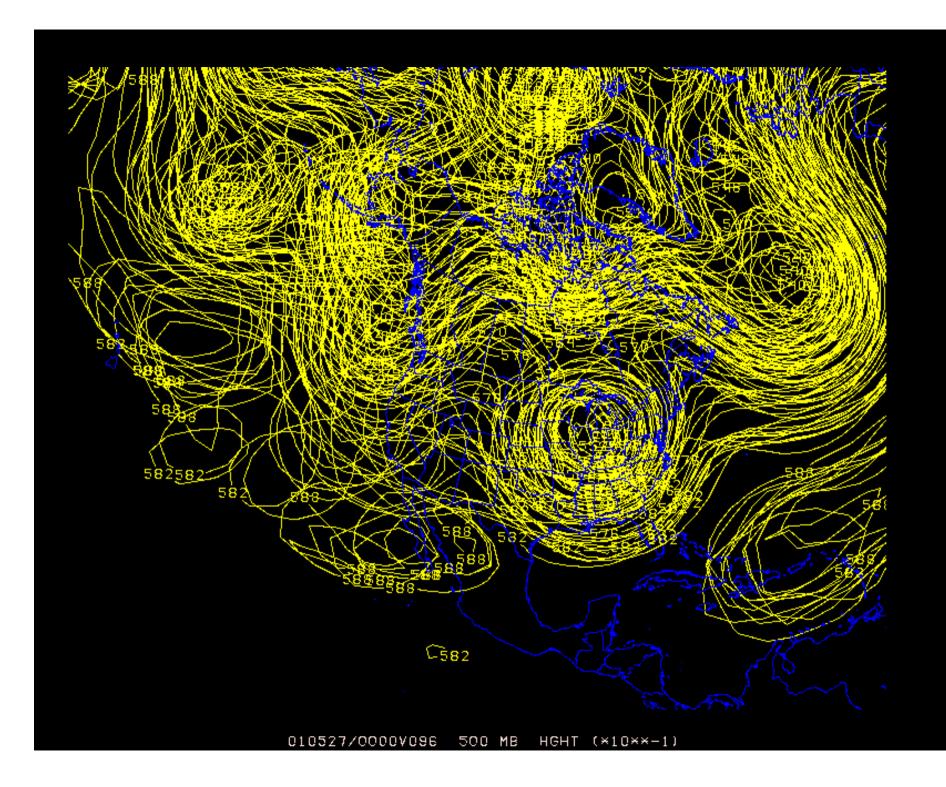


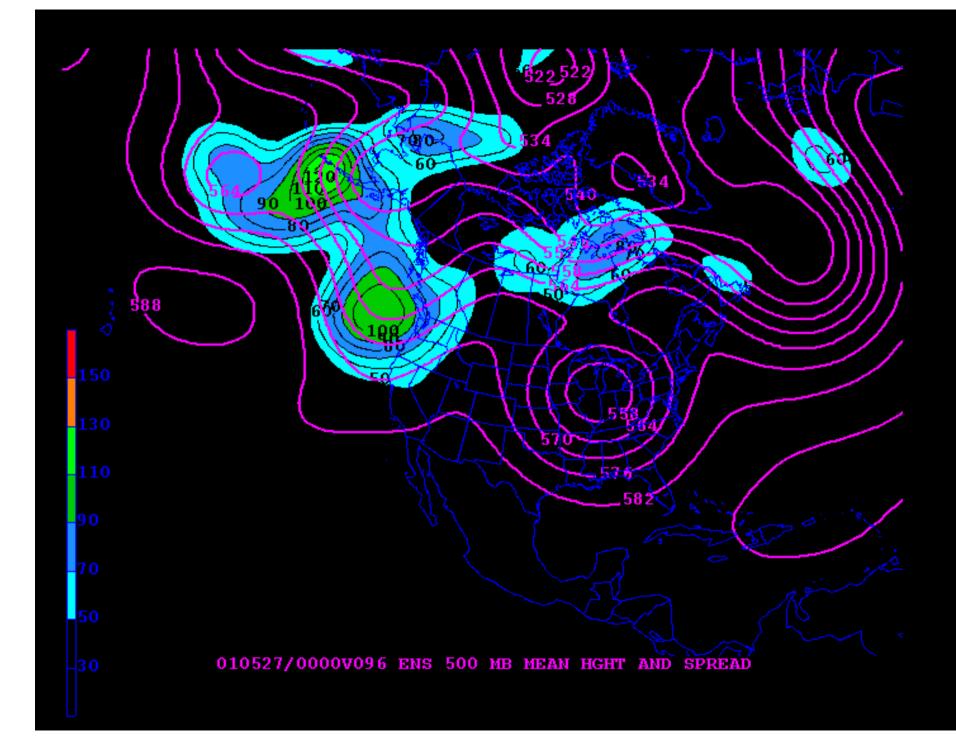
Ensemble Forecasts



Probability

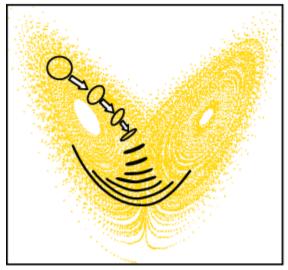


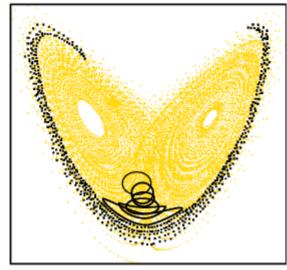




Lorenz Ensemble



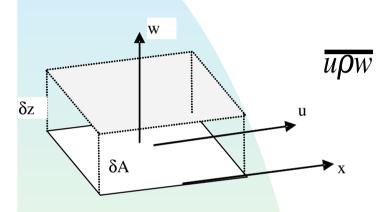




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Eddy flux measurements

